

1 The impact of NO_x emissions from lightning on the production 2 of aviation-induced ozone

3 **Arezoo Khodayari^{1,2}, Francis Vitt³, Daniel Phoenix¹ and Donald J. Wuebbles¹**

4 ¹ Department of Atmospheric Sciences, University of Illinois at Urbana-Champaign, USA.

5 ² Currently at the Department of Civil and Environmental Engineering, California State
6 University, Los Angeles, USA.

7 ³National Center for Atmospheric Research, Boulder, CO, USA.

8 *Corresponding author address: Arezoo Khodayari, Department of Atmospheric Sciences, University
9 of Illinois at Urbana-Champaign, 105 S. Gregory St., Urbana, IL, 61801. E-mail address:
10 akhoday2@illinois.edu

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Abstract

13 Due to the non-linear nature of ozone production in the troposphere, ozone production as a
14 function of aviation nitrogen oxide (NO_x = NO + NO₂) emissions varies based on the
15 background NO_x levels. Of the several different sources of background NO_x in the atmosphere,
16 NO_x from lightning (LNO_x) contributes a substantial amount of NO_x to the upper troposphere
17 and has an effect on the ozone production efficiency, even though the LNO_x source still has
18 significant uncertainty. In this study, CAM5, the atmospheric component of the Community
19 Earth System Model (CESM), was used to study the effect of uncertainties in NO_x emissions
20 from lightning on the production of aviation-induced ozone. Three sensitivity studies were
21 analyzed with varying LNO_x values of 3.7, 5, and 7.4 TgN/yr, representing the best current
22 range estimates for LNO_x. Results show a decrease in the aviation-induced ozone production
23 rate and radiative forcing (RF) as LNO_x increases. This is tied to the decreased ozone production
24 under NO_x saturated conditions. The ozone production per unit of NO_x emission from lightning
25 ranges from 2.38 TgO₃/TgN for the case with 3.7 TgN from lightning to 0.97 TgO₃/TgN for the
26 case with 7.4 TgN from lightning. Similarly, the O₃ RF decreases from 43.9 mW/m² for the 3.7
27 TgN/yr case to 34.3 mW/m² for 7.4 TgN/yr case. Understanding the current sensitivity of
28 aviation-induced ozone production to the LNO_x strength is important for reducing the
29 uncertainty in ozone production from aviation NO_x emissions.

30 1. **Introduction**

31 Nitrogen oxides (NO_x = NO + NO₂) are important to tropospheric chemistry due to their role
32 as a precursor of ozone [Crutzen, 1974]. NO_x emissions in the upper troposphere are particularly
33 important due to the pristine background atmosphere where additional NO_x emissions can have a
34 large impact on ozone production. A large anthropogenic source of NO_x emissions in the upper
35 troposphere is from commercial aircraft and these emissions are about four times more efficient
36 in producing ozone (O₃) than ground sources [Hodnebrog *et al.*, 2011].

37 Many past studies have examined the effects of aviation NO_x emissions on ozone [e.g.,
38 Brasseur *et al.*, 1998; IPCC, 1999; Grewe *et al.*, 2002; Köhler *et al.*, 2008; Hoor *et al.*, 2009;
39 Hodnebrog *et al.*, 2011; Myhre *et al.*, 2011; Köhler *et al.*, 2013; Jacobson *et al.*, 2013; Olsen *et al.*,
40 *et al.*, 2013b; Khodayari *et al.*, 2014]. Of these studies, there exists a considerable range in the
41 modeled difference in the increase in tropospheric ozone from aviation, from 2.3 to 3.0
42 TgO₃/TgN [Olsen *et al.*, 2013a] and 5.5 to 16.4 TgO₃/TgN [Lee *et al.*, 2010]. These differences
43 are likely due to several factors, among them, the upper tropospheric budget of background NO_x
44 and the non-linearity of ozone production relative to the background NO_x concentration [IPCC,
45 1999]. Two dominant sources of NO_x in the upper troposphere are production from lightning
46 (LNO_x) and emissions from aviation, with lightning being the largest non-aviation source of
47 NO_x [WMO, 1999]. However, the magnitude of NO_x emissions from lightning has been one of
48 the least known in the global NO_x budget, with a current uncertainty range of 1-20 with the most
49 probable range being 2-8 TgN/yr [Schumann and Huntreiser, 2007, and references therein].
50 Additionally, with the large error in the retrieval process of observational data of cloud to ground
51 flashes and the uncertainty of energy dissipated from intracloud flashes, lightning
52 parameterizations in CTMs may not accurately represent the vertical distribution of LNO_x [Ott
53 *et al.*, 2010; Miyazaki *et al.*, 2014]. While the contribution from aviation is reasonably well
54 known [Wilkerson *et al.*, 2010; Olsen *et al.*, 2013a], the contribution from lightning is harder to
55 accurately represent.

56 Another complication is that the production of ozone as a function of the background
57 NO_x level is non-linear [Isaksen *et al.*, 1978; Lin *et al.*, 1988; Grooß *et al.*, 1998; Jaeglé *et al.*,
58 1999]. Ozone production is sensitive to the background concentrations of NO_x and nonmethane

59 hydrocarbons (NMHC or volatile organic compounds, VOC). Therefore, the efficiency and sign
60 of ozone production depends on the background atmosphere and therefore, LNO_x strength.

61 Numerous studies have shown that aviation NO_x emissions increase ozone production in the
62 upper troposphere [e.g. *Brasseur et al.*, 1998; *IPCC*, 1999; *Grewe et al.*, 2002; *Köhler et al.*,
63 2008; *Hoor et al.*, 2009; *Hodnebrog et al.*, 2011; *Myhre et al.*, 2011; *Köhler et al.*, 2013;
64 *Jacobson et al.*, 2013; *Olsen et al.*, 2013b; *Khodayari et al.*, 2014]. Other studies have also
65 shown that increases in the LNO_x source increase ozone production in the background
66 atmosphere [*Lawrence et al.*, 1995; *Brasseur et al.*, 1996; *Lee et al.*, 1997; *Berntsen and Isaksen*,
67 1999; *Kraabøl et al.*, 2002; *Labrador et al.*, 2005; *Schumann and Huntreiser*, 2007; *Wild*, 2007;
68 *Allen et al.*, 2010; *Liaskos et al.*, 2015]. A few studies have shown that increases in the LNO_x
69 source decrease the aviation-induced ozone production efficiency due to the high NO_x/low VOC
70 atmosphere [*Brasseur et al.*, 1996; *Berntsen and Isaksen*, 1999]. Brasseur et al. [1996] found the
71 production of aviation NO_x-induced ozone was reduced by a factor of 2 when the lightning
72 source was increased from 5 TgN/yr to 10 TgN/yr at 10 km. Berntsen and Isaksen [1999] found a
73 1.5- 2.5 ppb maximum reduction of aviation-induced ozone when the NO_x source from lightning
74 was increased from 5 TgN/yr to 12 TgN/yr.

75 Given the improvements in understanding of atmospheric chemistry and physics since these
76 earlier studies, we found it worthwhile to investigate the effect of LNO_x on the chemistry of the
77 background atmosphere and on the production of aviation-induced ozone . Our study investigates
78 the aviation-induced O₃ production and radiative forcing from O₃ for three different LNO_x
79 magnitudes using the most up-to-date aviation NO_x emission inventory and a state-of-art
80 climate-chemistry model. The updates in the emission inventory and the improvements in
81 climate-chemistry modeling have led to higher model resolution, a more complete chemical
82 mechanism and a more detailed representation of aerosols processes that are central to ozone
83 chemistry and should lead to a more accurate estimate of the impact of LNO_x on the production
84 of aviation-induced ozone.

85 This study will investigate the effects of lightning NO_x on ozone production with aviation
86 NO_x emissions using version 5 of the Community Atmosphere Model (CAM5) which is the
87 atmospheric component of Community Earth System Model (CESM). Three sets of simulations
88 were done with varying LNO_x values of 3.7, 5.0, and 7.4 TgN/yr, in accordance with the current

89 best estimate for LNO_x, to analyze the effect of LNO_x on the production of aviation-induced
90 ozone. The rest of the paper is organized as follows. Section 2 gives a description of the model
91 used and the simulations done, section 3 presents the results and discussion, and section 4
92 provides the concluding remarks.

93 2. **Model Description and Simulation Setup**

94 The Community Atmosphere Model version 5 (CAM5) was used for this study. CAM5 was
95 configured with a horizontal resolution of 1.9° latitude x 2.5° longitude and 56 vertical levels
96 from the surface to ~3.5 hPa. The vertical resolution in the vicinity of the tropopause was about
97 1.3 km. The photolysis scheme and gas phase mechanism, which includes 133 species and 330
98 photochemical reactions, is the same as in CAM4 and has been detailed in *Lamarque et al.*
99 [2012]. The main improvement in CAM5 over its predecessor is its modal aerosol module
100 (MAM) [*Liu et al.*, 2012]. Two versions of MAM were developed, one with three lognormal
101 modes (MAM3) and one with seven lognormal modes (MAM7). For this study we used MAM7,
102 which represents Aitken, accumulation primary carbon, fine dust and sea salt, and coarse dust
103 and sea salt modes. The mass and number mixing ratio for all aerosols in each mode is
104 calculated, and the size distribution of aerosols, internal and external mixing, and physical,
105 chemical and optical properties are simulated [*Liu et al.*, 2012].

106 Simulations were run with specified dynamics (“off-line” mode) to reduce the year-to-year
107 climate variability and better detect the aviation NO_x signal. Being run in “off-line” mode,
108 changes in the chemical constituents do not affect the dynamics. GEOS DAS v5.1 meteorology
109 for the year 2005 [*Rienecker et al.* 2008] was used since it was the closest available assimilated
110 meteorology data to the year of interest (2006). The AEDT aviation emissions analyses were
111 used to obtain the aviation emissions for the year 2006 [*Wilkerson et al.*, 2010; *Olsen et al.*,
112 2013a; *Brasseur et al.*, 2013]. The aviation NO_x emissions had a horizontal resolution of 1°
113 latitude x 1° longitude, vertical resolution of 150 m, and hourly temporal resolution. The total
114 NO_x emissions amounted to 0.82 TgN/yr. The IPCC RCP4.5 scenario [*van Vuuren et al.*, 2011]
115 for year 2005 was used to obtain the background emissions of non-aviation short-lived species
116 (e.g., NO_x, VOCs) while the Model of Emissions of Gases and Aerosols from Nature (MEGAN)
117 [*Guenther et al.*, 2006] was used to obtain the background emissions of biogenic species.
118 Longer-lived species, e.g., carbon dioxide (CO₂), methane (CH₄), chlorofluorocarbons (CFCs),

119 and nitrous oxide (N₂O), were specified as boundary conditions based on the IPCC RCP4.5
120 scenario.

121 To investigate aviation-induced changes in O₃, three sets of two simulations were performed;
122 a simulation with aviation NO_x emissions and a simulation without aviation NO_x emissions (i.e.
123 control simulation), each set with a different LNO_x level. The simulations were run for 7 years to
124 reach steady-state while cycling through the 2005 meteorology field to remove the year-to-year
125 variability. Data from the 7th year was used for analysis. Aviation-induced changes in O₃ were
126 calculated as the difference between the two runs. LNO_x values of 3.7, 5, and 7.4 TgN/yr were
127 chosen in accordance with the range of current estimates, where the best estimate is 5 TgN/yr
128 [Schumann and Huntrieser, 2007]. The set with 3.7 Tg N/yr will be referred to as the reference
129 case and the set with 5 TgN/yr and 7.4 TgN/yr will be referred to as case 2 and case 3,
130 respectively. Since the only difference between the three sets of simulations were the LNO_x
131 levels, the differences in the aviation-induced O₃ changes between the three sets were solely due
132 to the differences in the level of LNO_x.

133

134 3. Results and Discussion

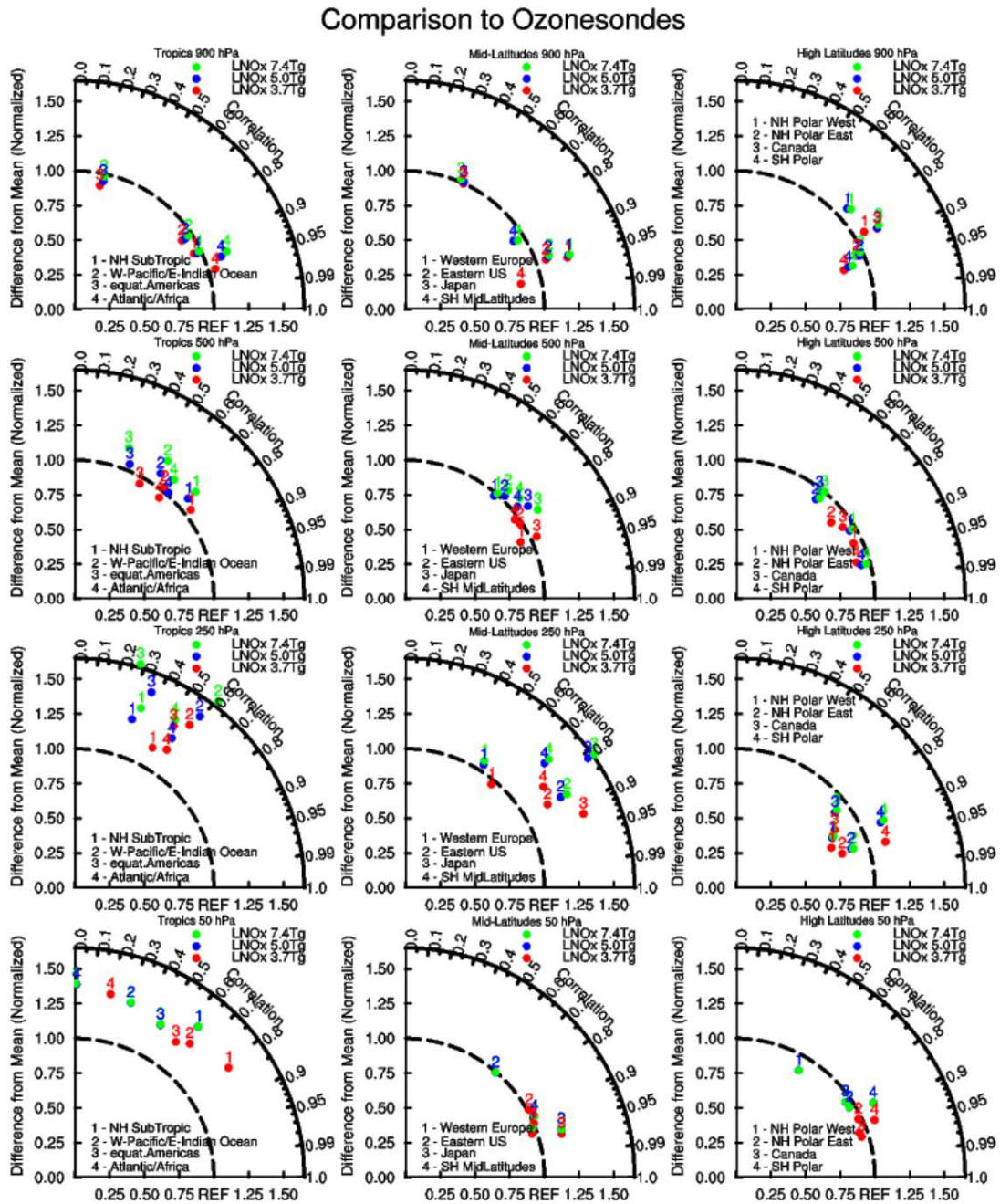
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136 3.1. Chemistry diagnosis of the background atmosphere

137 It has been previously shown that CAM reasonably reproduces the effects of aviation NO_x-
138 induced emissions on the distribution of tropospheric O₃ and NO_x [Weber, 2011 and Olsen et al.,
139 2013b]. However, because of the central role of ozone in the troposphere and stratosphere and its
140 relation to NO_x budget, simulated ozone under different level of NO_x emissions from
141 lightning (LNO_x) and at representative altitudes is evaluated using an ozonesonde climatology
142 [Tilmes et al., 2012]. This climatology comprises observations for the years 1995-2011 and
143 includes averaged ozone profiles for 41 different ozonesonde stations that are grouped into 12
144 regions. For the purpose of our study, we evaluate ozone at four pressure levels covering the
145 troposphere and lower stratosphere (50, 250, 500, and 900 hPa) over the 12 areas that are
146 grouped into three larger regions (Tropics, Mid-Latitudes, and High Latitudes), as shown in
147 Figure 1. Model results are interpolated horizontally to all the stations within each region, and

148 averaged over each region. Figure 1 shows the comparison between model and observation in
 149 Taylor-like diagrams for each of the corresponding pressure levels and regions.

150



152 Figure 1. Taylor diagram of modeled annual average background ozone from the control runs
 153 against ozonesonde climatology for four pressure levels and three latitudinal regions. REF along
 154 the abscissa denotes the observations while the radial distance describes the normalized bias. The

155 correlation between modeled and observed seasonality is shown by the azimuthal angle. It is
156 noted that most aircrafts fly around 250 hPa.

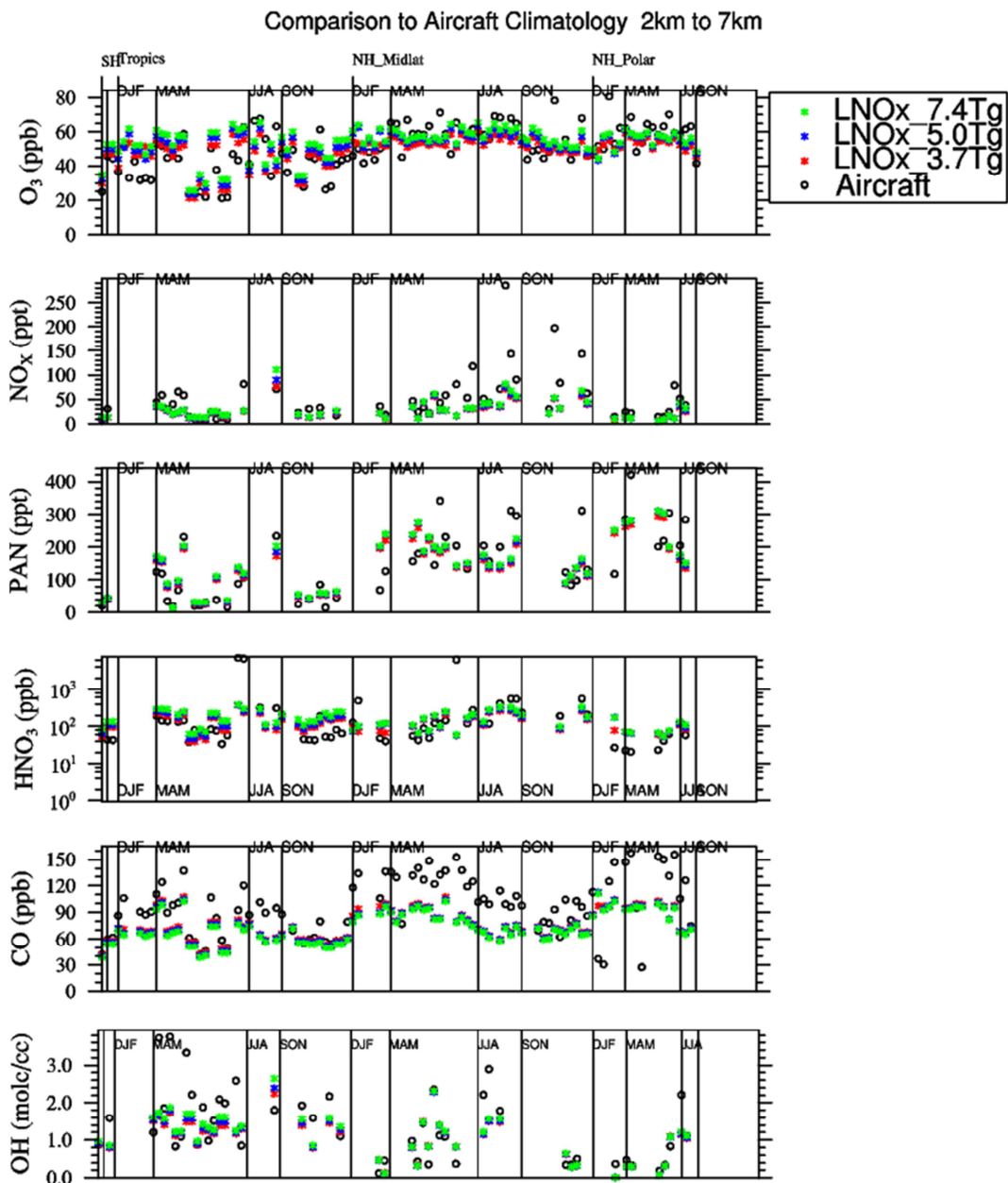
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158 As shown in Figure 1, the simulated ozone is within 10% of the observation in the boundary
159 layer (900 hPa) and at 500 hPa which is the range of uncertainty for observation and is a good
160 agreement. Moreover, with an exception of few locations in tropics the seasonal cycle is well
161 reproduced at the boundary layer (900 hPa) and 500 hPa. The bias in the tropics is likely due to a
162 model estimated tropopause that is lower than observed and likely too much transport of ozone
163 into the troposphere [Lamarque *et al.*, 2012]. It is noted that the simulated ozone is very similar
164 under different level of LNO_x in the boundary layer (900 hPa) and 500 hPa. On the other hand,
165 the simulated ozone is less in agreement with observation at higher altitudes (e. g. at 250 and 50
166 hPa), especially in mid-latitudes and Tropics. Basically, at these altitudes the simulated ozone is
167 mainly overestimated by about 25% in mid-latitudes and 50% in Tropics. The seasonal cycle
168 with the exception of in Tropics is relatively well produced. Overall, as illustrated in Figure 1,
169 ozone is well simulated in the troposphere, in the upper troposphere-lower stratosphere (UTLS)
170 and stratosphere and it is overestimated in the tropical transition layer. It is also noted despite
171 what was the case at the boundary layer and in the troposphere, at UTLS the simulated ozone
172 varies at different level of NO_x emissions from lightning and is lower at lower level of LNO_x.
173 As shown in Figure 1 at UTLS the simulated ozone is in a better agreement with observation at
174 lower level of LNO_x (i. e. 3.7 Tg N). To further investigate the effect of different level of NO_x
175 emissions from lightning on the chemistry of the background atmosphere, the simulated level of
176 O₃, NO_x, PAN, HNO₃, CO as well as OH, are compared to aircraft observations between 2-7 km,
177 where the majority of the observations were taken [Emmons *et al.*, 2000] as illustrated in Figure
178 2.

179 In comparison to aircraft data, the simulated ozone is higher under higher level of NO_x
180 emissions from lightning, especially in the tropics and mid-latitudes, in agreement with
181 ozonesonde observations. The simulated NO_x is also in a good agreement with available aircraft
182 observations, but it is underestimated in summer and fall in NH high-latitudes. The simulated
183 NO_x is also similar under different level of NO_x emissions from lightning which indicates a
184 higher shift of additional NO_x to other NO_y species under a higher level of NO_x emission from

185 lightning as shown in Figure 2. This observation is in agreement with higher level of simulated
 186 HNO₃ and PAN, and OH under higher level of LNO_x. The simulated CO in line with higher
 187 production of ozone is slightly lower under higher level of NO_x emissions from lightning and is
 188 underestimated in all simulations. Overall, the differences between simulated results under
 189 different level of NO_x emissions from lightning are within the variability of the observations.

190



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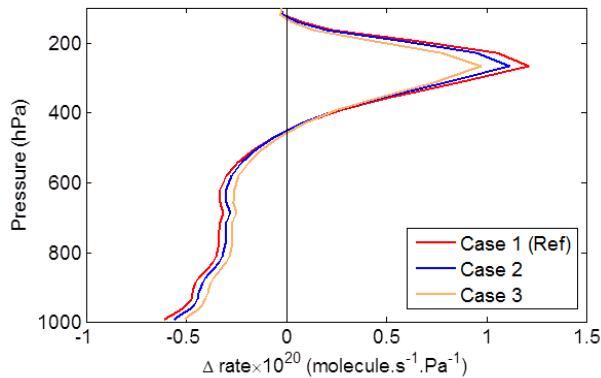
192 Figure 2. Comparison between aircraft observations over different regions and seasons and
193 different model simulations derived by different level of LNO_x emissions, averaged between 2
194 and 7 km, for ozone, NO_x, PAN, HNO₃, CO, and OH based on an updated version of the aircraft
195 climatology by [Emmons *et al.*, 2000], as described in detail in [Tilmes *et al.*, 2015]. As
196 discussed in Emmons *et al.* (2000), for each aircraft campaign, regions with high frequency
197 occurrence of vertical profiles from the aircraft were recognized. Mean profiles of the examined
198 species across all the campaigns were compiled over these regions with a 1 km vertical
199 resolution. Profiles that were outliers were removed. The multiple data points within each region
200 and season represent different location at which observations were made and model data were
201 interpolated to the locations of the observational data for comparison purposes.

202

203 3.2. Aviation-induced changes in ozone

204 Due to complex photochemistry of ozone production, additional units of NO_x emissions could
205 result in both production and destruction of ozone [Lin *et al.*, 1988]. As such, the production of
206 ozone does not follow linearly with increasing concentrations of NO_x since the destruction rates
207 of NO_x and odd hydrogen depend on the initial concentration of local NMHC and NO_x in the
208 background atmosphere [Lin *et al.*, 1988]. Since NO_x is a catalyst in ozone production process,
209 its increase could increase the background O₃ levels if the initial concentration of local NMHC
210 and NO_x in the background atmosphere favors ozone production (i.e. low NO_x/NMHC ratio).
211 According to our results and as shown in Figures 1 and 2 the additional LNO_x emissions led to
212 more ozone production in the background atmosphere, such that there is a maximum zonal O₃
213 increase of 63% between case 3 (7.4 TgN/yr) and the reference case (3.7 Tg N/yr) between 100-
214 200 hPa over 30-60° N. However, due to high NO_x/NMHC ratio under high LNO_x emissions,
215 aviation-induced production of ozone at cruise altitudes decreased with increasing LNO_x
216 emissions. Figure 3 shows the changes in aviation-induced ozone production in the free
217 troposphere for the 3 cases. The production rate of ozone has been defined as the instantaneous
218 production minus loss for ozone (molecules cm⁻³ S⁻¹) for a model grid box integrated over the
219 global and expressed in Pa⁻¹ to account for the thickness of the box. As shown in Fig. 3, the net
220 aviation-induced ozone production at cruise altitude is largest in the reference case and smallest
221 for case 3. Therefore, aviation-induced ozone production rate is greatest for the case with the
222 lowest LNO_x level.

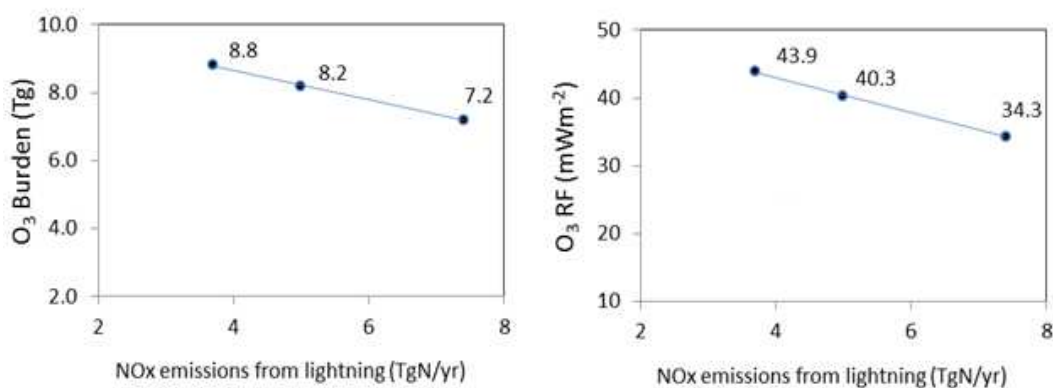
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224

225 Figure 3. Aviation-induced ozone net production rate in the free troposphere for the 3 cases. The
 226 net production rate of ozone for the reference case (case with 3.7 Tg N/yr NO_x emissions from
 227 lightning) is indicated by the red line. The blue and orange lines show the net production rate of
 228 ozone for case 2 and case 3, respectively.

229 The global aviation-induced O₃ burden and O₃-RF are shown in Figure 4 and Table 1. As
 230 expected by the higher aviation-induced ozone production rate under the reference case (low
 231 LNO_x) the aviation-induced global tropospheric O₃ burden and O₃-RF are largest for the
 232 reference case with a global O₃ burden of 8.8 Tg and O₃-RF of 43.9 mWm⁻². For a doubling of
 233 the LNO_x source from 3.7 to 7.4 TgN/yr, the aviation induced global tropospheric O₃ burden and
 234 O₃-RF decrease by 18% and 22%, respectively. Furthermore, the global tropospheric aviation-
 235 induced ozone burden per Tg of NO_x emitted from lightning decreases from 2.38 TgO₃/TgNO_x
 236 for the reference case to 0.97 TgO₃/TgLNO_x for case 3, as shown in Table 1.



237

238 Figure 4. Aviation-induced O₃ RF (mW/m²) and O₃ Burden (Tg) from 3 different LNO_x sources
 239 (3.7 Tg, 5.0 Tg, and 7.4 Tg).

240 Table 1. LNO_x source and corresponding aviation-induced O₃ burden, O₃ Radiative Forcing, and O₃
 241 burden per Tg of LNO_x

LNO _x Source (TgN)	O ₃ burden (Tg)	O ₃ RF (mW/m ²)	O ₃ /LNO _x (Tg O ₃ /TgN)
3.7	8.8	43.9	2.38
5.0	8.2	40.3	1.64
7.4	7.2	34.3	0.97

242

243 In previous studies, *Brasseur et al.* [1996] and *Berntsen and Isaksen* [1999] also looked at the
 244 effect of varying LNO_x levels on aviation-induced ozone at cruise altitudes. *Brasseur et al.*
 245 [1996] found that for a doubling of the LNO_x source from 5 TgN/yr to 10 TgN/yr, the zonally
 246 averaged aviation-induced tropospheric ozone production was reduced by a factor of 2, from
 247 4.5% to 2.4% of the background ozone in the northern hemisphere. Our results show an 14%
 248 decrease in the northern hemisphere (or a reduction by a factor of 1.2, from 3.4% to 2.8%) in the
 249 zonally averaged aviation-induced tropospheric ozone production for a doubling of the LNO_x
 250 source from 3.7 to 7.4 TgN/yr. *Berntsen and Isaksen* [1999] found, for an increase of the LNO_x
 251 strength from 5 TgN/yr to 12 TgN/yr, a maximum reduction of aviation-induced ozone
 252 production of 1.5-2.5 ppb in northern mid- and high-latitudes during the summer. In our study,
 253 we estimate a maximum increase of 1 ppb in aviation-induced ozone production from a decrease
 254 in the LNO_x strength from 3.7 TgN/yr to 7.4 TgN/yr over northern mid- and high-latitudes
 255 during summer.

256 The difference between these two studies and ours is likely due to the non-linear relationship
 257 between the amount of NO_x emissions from lightning and the amount of aviation-induced ozone
 258 as well as the following reasons. The 3-dimensional chemical tracer model (CTM) used by
 259 *Brasseur et al.* [1996] had a 5 x 5 degree horizontal resolution with 25 vertical levels and
 260 included 50 species, 6 NMHCs and several oxygenated organic species including PAN and other
 261 organic nitrates. In the study by *Berntsen and Isaksen* [1999], the 3-dimensional CTM used had
 262 an 8 x 10 degree horizontal resolution with 9 vertical layers. The model had photochemical
 263 scheme of 55 chemical components and 120 gas-phase reactions. The chemical scheme was
 264 extended to include the degradation of propane, which gives acetone. However, there were no

265 direct emissions of acetone (an important source of HO_x) in the model. This may have important
266 implications on their result as *Müller and Brasseur* [1999] found a 20% reduction in O₃ mixing
267 ratios when emissions of acetone were not included. Our study uses CAM5 with chemistry, on a
268 1.9 x 2.5 degree horizontal resolution with 30 vertical levels. CAM5 has tropospheric and
269 stratospheric chemistry with 133 species, 330 photochemical reactions, and a modal aerosol
270 module. We also use aviation NO_x emissions from the year 2006 (i. e. 0.82 TgN) while *Brasseur*
271 *et al.* [1996] and *Berntsen and Isaksen* [1999] used aircraft emissions from 1990 (i. e. 0.44 TgN)
272 and 1996 (i. e. 0.52 TgN), respectively. The updates in our model resolution and chemistry
273 scheme, with more recent aviation emissions, provide a current evaluation of the impact of NO_x
274 emissions from lightning on the production of aviation-induced O₃.

275 As improvements in chemistry-climate models are important for evaluating aviation-induced
276 ozone production, reducing the LNO_x uncertainty is also important. As was seen in Figure 1 and
277 2 the level of NO_x emission from lightning affects the simulation of the background atmosphere
278 and its agreement with observations. In this study the simulated ozone in the background
279 atmosphere was in a better agreement with observation under a low LNO_x emission case (3.7 Tg
280 N/yr). *Schumann and Huntrieser* [2007] estimated a range of 2-8 TgNyr⁻¹ with a best estimate of
281 5 TgNyr⁻¹ in a review of 30 years of lightning studies. The techniques these studies used to
282 estimate LNO_x fell into three categories: flash extrapolation, storm extrapolation, and global
283 model fit. Of these methods, the global model fit had the lowest range of uncertainty. However,
284 at the time of publication, only a subset of observations had been compared to the model
285 simulations. Furthermore, the model estimates were based on meteorological parameters and did
286 not account for all the parameters that are important for lightning and LNO_x production
287 [*Schumann and Huntrieser*, 2007].

288
289 More recent studies have found similar best estimates with reduced uncertainty. *Murray et al.*
290 [2012] estimated a global annual LNO_x source of 5.5-6.5 TgNyr⁻¹ after applying a scaling factor
291 to improve the representation of lightning in CTMs, using Lightning Imaging Sensor (LIS)
292 /Optical Transient Detector (OTD) satellite observations. *Stravroukou et al.* [2013] estimated a
293 range of 3.3-5.9 TgNyr⁻¹ after considering recent developments in the understanding of NO_x sink
294 pathways with top-down emissions derived from inverse modelling. *Miyazaki et al.* [2014]
295 estimated a global annual LNO_x source of 4.9-7.7 TgNyr⁻¹ from assimilating satellite

296 observations of NO₂, O₃, HNO₃, and CO in a chemical transport model. While more recent
297 studies have better constrained the LNO_x source, improvements still need to be made in the
298 simulated lightning variability, the accuracy of simulated atmospheric electricity, convective
299 parameterizations, and vertical emissions profiles [*Schumann and Huntrieser, 2007*].
300

301 4. Conclusions

302 We used the most up-to-date aviation emission inventory and a state-of-the art climate-
303 chemistry model, CAM5, to evaluate the impact of LNO_x on the production of aviation-induced
304 ozone. CAM5 has a higher model resolution, more complete chemical mechanism and a more
305 detailed representation of aerosol processes compared to previous models that were used for
306 evaluating the impact of LNO_x on the production of aviation-induced ozone. As such the result
307 presented in this study can be considered the most current evaluation of the impact of LNO_x on
308 the production of aviation-induced ozone.

309 CAM5 simulations with varying LNO_x source strengths show that increasing the LNO_x
310 source from 3.7 to 7.4 TgN/yr decreases the global mean aviation-induced O₃ burden and O₃-RF
311 by about 18% and 22%, respectively with a maximum reduction in the aviation-induced surface
312 ozone perturbation of 1 ppb. Additionally, for a doubling of the LNO_x strength, the NH zonal
313 mean O₃ perturbation is reduced by a factor of 1.2. The greatest change in the net ozone
314 production rate in the three cases is in the Northern Hemisphere at subsonic aircraft cruise
315 altitudes. Previous studies show similar relationships between LNO_x levels and aviation-induced
316 ozone production, but the magnitude of these changes is lower in our study.

317 Our study provides a current estimate of the effects of LNO_x on aviation-induced ozone. The
318 range in uncertainty for both aviation-induced tropospheric ozone and the NO_x source from
319 lightning is large (2.3 to 3.0 TgO₃/TgN [*Olsen et al., 2013a*] and 5.5 to 16.4 TgO₃/TgN [*Lee et*
320 *al., 2010*] for aviation-induced tropospheric ozone and 2 to 8 TgN [*Schumann and Huntrieser,*
321 *2007*] for LNO_x). The uncertainty in aviation-induced tropospheric ozone is likely due to
322 differences in the details representations of the physics and chemistry of the background

323 atmosphere in different models among which representation of NO_x emissions from lightning is
324 an important one [Olsen *et al.*, 2013a]. As such assessing the current model sensitivity of
325 aviation-induced ozone to LNO_x is important to reduce the uncertainty in aviation-induced
326 ozone production. More importantly, a better understanding of the magnitude of the NO_x
327 sources and sinks from lightning is needed to better quantify the impact of aviation emissions on
328 ozone production. The results from this study suggest that there remain uncertainties in the
329 effects of NO_x emissions from lightning on aviation-induced ozone. It further points to the
330 necessity of more detailed critical testing of models sensitivity to the effect of NO_x emissions
331 from lightning and a better understanding of the magnitude of the NO_x emission from lightning
332 as perhaps the most important path to reducing these uncertainties.

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We calculated the impact of NO_x emissions from lightning on the production of aviation-induced ozone.

Three sensitivity studies were conducted with varying levels of NO_x emissions (LNO_x) from lightning.

By doubling LNO_x the global mean aviation-induced O₃ burden decreased by about 18%.

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